CHANDLER WOBBLE VARIABILITY

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Precession, nutation and polar motion – astronomy or geophysics?

Precession and Nutation 岁差和章动 of the Earth axis 26 000, 18.6 years, 1, 0.5 year caused by Sun and Moon tides



Earth's pole motion 极点运动 up to 10 meters caused by geophysical effects – momentum exchange between the Ocean, Atmosphere, and Solid Earth



Polar motion trajectory



Polar motion principal components trend, annual and Chandler (433 days) wobble

8. C. Canster



Polar motion spectrum



Variable Chandler and Annual Wobbles in Earth's Polar Motion During 1900–2015

Guocheng Wang¹ · Lintao Liu¹ · Xiaoqing Su² · Xinghui Liang¹ · Haoming Yan¹ · Yi Tu³ · Zhonghua Li¹ · Wenping Li⁴ Surv Geophys (2016) 37:1075–1093 DOI 10.1007/s10712-016-9384-0



Fig. 8 Reconstruction of the CW using the FBPBPF method

Amplitude and phase variations of Earth's Chandler wobble under continual excitation

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Fig. 1. (a) The *x* and *y* components of the Hipparcos polar motion series (for the period 1899.7–1992.0) concatenated with EOP-CO1 (for the period 1992.0–2010.0) for the total span of 110 years at 5-day intervals. (b) The Chandler wobble series m(t) obtained from (a) after removing the least-squares estimates of the annual wobble and a linear trend.

Chandler wobble: two more large phase jumps revealed

Zinovy Malkin and Natalia Miller



Panteleev's filtering in the Chandler band



Filtered Chandler wobble and its envelope



CW amplitude and phase model and prediction



Comparison of left and right parts of Euler-Liouville equation for polar motion



p(t) – complex polar motion trajectory

 $\chi(t)$ – excitation

$$\sigma_c = \sigma_e \frac{\left(1 - \frac{k_2}{k_s}\right)}{\left(1 + e\frac{\tilde{k}_2}{k_s}\right)} = 2\pi f_c \left(1 + \frac{i}{2Q}\right).$$



Splitted Chandler wobble spectral line



Modelling amplitude changes in the Chandler Wobble excitation





How to select filter width?



Investigation of inverse problem solution uncertainty

$$p(t) = A\chi(t) + \delta$$

$$\breve{\chi}(t)$$
=R $p(t)$

direct problem, δ – observational noise

solution of the inverse problem, *R* - regularizing algorithm

$$\breve{\chi} = argmin \|A\breve{\chi} - p\|$$

there is error *h* in operator *A* which depends on knowledge of CW frequency f_c and quality factor Q.

 $h \| \breve{\chi} \|$

- error in *p* related to operator uncertainty

R also depends on CW frequency f_c and quality factor Q, but the main parameter of the regularizing algorithm is the filter width parameter f_0

the only way to access the solution $\breve{\chi}$ uncertainty is to provide the diameter of the set, it is selected from, consistened with uncertainties *h*, δ and a priory conditions, used in *R*



filter width parameter $f_0 = 1/T_0$

Comparison of exact modeled excitation with what was reconstructed through corrective filtering (inversion+filter) for different values of filter width parameter $f_0=1/T_0$





The discrepancy valued $||A\chi - p||$ depending on the filter parameter selected



 $\mu = ||A\breve{\chi} - p|| + \delta + h||\breve{\chi}||$ - generalized uncertainty

This proves that the optimal half-width parameter of the filter $f_0=1/T_0$ is about 1/20 - 1/25 (in our previous work $f_0=1/25$ year ⁻¹)

Angular momentum – geophysical excitations

Mass component

Motion component

$$\begin{split} \mathbf{H}(t) &= \mathbf{H}^{mass}(t) + \mathbf{H}^{motion}(t) = \int \rho(r, t) \mathbf{r} \times [\mathbf{\Omega} \times \mathbf{r} + \mathbf{v}(r, t)] dV. \\ \mathbf{H}_x^{mass}(t) &= -\Omega \frac{r^4}{g} \int \int p(\phi, \lambda) \sin \phi \cos^2 \phi \cos \lambda d\phi d\lambda, \qquad \mathbf{H}_x^{motion}(t) = \frac{r^3}{g} \int \int \int [\cos \phi \sin \lambda v(r, t) - \sin \phi \cos \phi \cos \lambda u(r, t)] d\phi d\lambda, \\ \mathbf{H}_y^{mass}(t) &= -\Omega \frac{r^4}{g} \int \int p(\phi, \lambda) \sin \phi \cos^2 \phi \sin \lambda d\phi d\lambda, \qquad \mathbf{H}_y^{motion}(t) = \frac{r^3}{g} \int \int \int [-\cos \phi \cos \lambda v(r, t) - \sin \phi \cos \phi \sin \lambda u(r, t)] dp d\phi d\lambda, \\ \mathbf{H}_z^{mass}(t) &= \Omega \frac{r^4}{g} \int \int p(\phi, \lambda) \cos^3 \phi d\phi d\lambda, \qquad \mathbf{H}_z^{motion}(t) = \frac{r^3}{g} \int \int \int \cos^2 \phi u(r, t) dp d\phi d\lambda, \end{split}$$

Effective angular momentum

 $\chi_{x,y}^{motion} = \frac{1.5913}{\Omega(C-A)} \mathbf{H}_{x,y}^{motion}, \qquad \chi_z^{motion} = \frac{0.998}{\Omega C} \mathbf{H}_z^{motion},$ $\chi_{x,y}^{mass} = \frac{1.098}{\Omega(C-A)} \mathbf{H}_{x,y}^{mass}, \qquad \chi_z^{mass} = \frac{0.753}{\Omega C} \mathbf{H}_z^{mass}.$









Chandler wobble, comparison of excitations



Chandler wobble, comparison of excitations



Planteleev's filtering in prograde and retrograde Chandler band



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CHANDLER EXCITATION

Geodetic and geophysical excitations in prograde and retrograde Chandler bands



Conclusions

- Chandler wobble can be extracted by different types of filtering
- CW amplitude and phase changes observed in our epoch and over more than 100 years can be crucial for Chandler wobble understanding
- spectral line of CW is splitted, what manifests its variability in time, may be coupled oscillations in the Earth systems are responsible for regularities
- if 80 and 40-years quasi-periodic changes in CW amplitude are real, they would present in excitation (assuming stability of equation parameters)
- modelling proves, that ~40-year changes in CW amplitude are related to the repetitions of ~20-year excitation and ~20-year damping epochs
- the so-called "free" Chandler mode actually is not free, being provided by atmospheric and oceanic excitations with, possibly, hidden regularity

Thank you for attention!

