



Excitation of the Earth's Chandler wobble by the North Atlantic double-gyre

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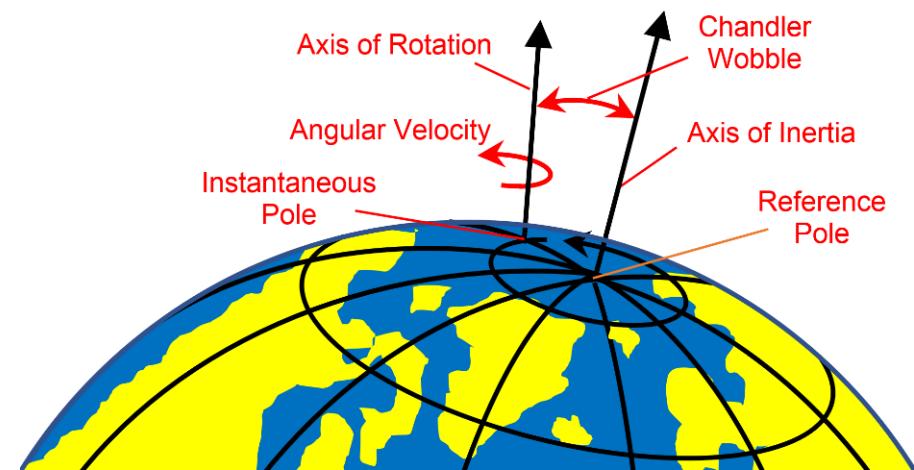
Chandler wobble

Chandler Wobble, the main component of polar motion, is a 14-month free motion, the period of which is determined by elliptic geometry and the rigidity of the Earth .

Chandler wobble
equations

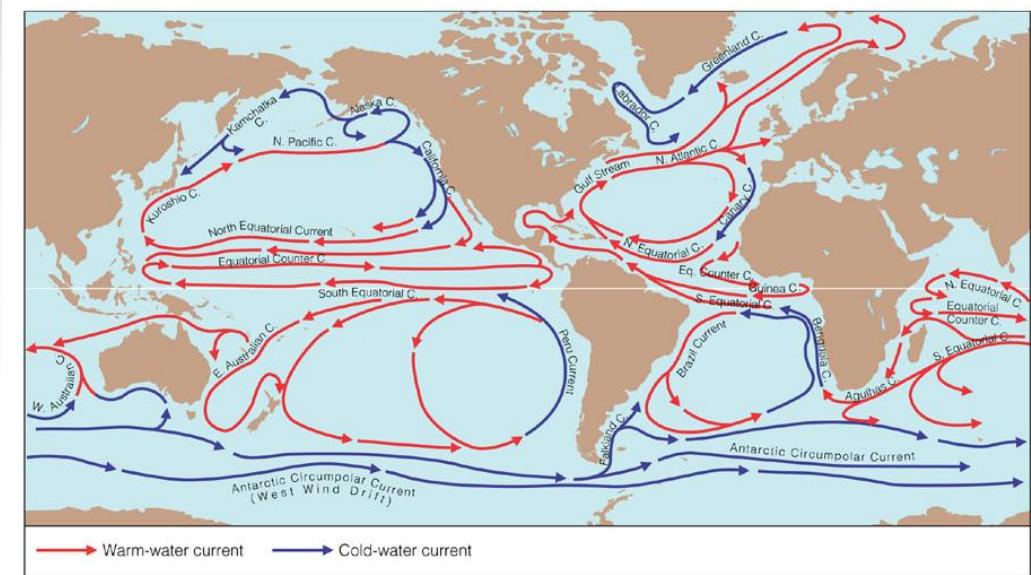
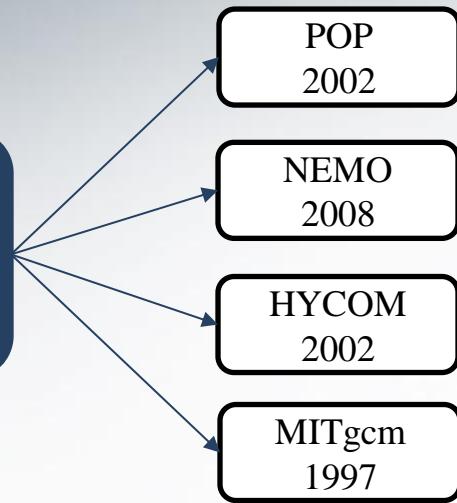
$$\frac{i}{\sigma_0 + i / 2Q} \frac{d\mathbf{m}}{dt} + \mathbf{m} = \boldsymbol{\psi} = \left[1 - \frac{i}{\Omega} \frac{d}{dt} \right] \{ \alpha \mathbf{c} + \beta \mathbf{h} \}$$

$$T_0 = \frac{2\pi}{\sigma_0} \approx 433 \text{ days}$$



Background: general circulation ocean models & double-gyre models

General
Circulation
Models



double-gyre models

Munk 1950

Holland
1978

Shen et al.
1999

Berloff 2005

Karabasov
et al. 2009

Maddison et
al. 2015

Shevchenko &
Berloff 2015



Approach

HYCOM (HYbrid Coordinate Ocean Model):

A general ocean model with realistic continent boundaries, bottom topography and time varying wind forcing

Quasi-geostrophic model (QG):

A mid-latitude approximation of the double-gyre problem with steady wind forcing, flat bottom topography and longitudinal-latitudinal boundary walls



HYbrid Coordinate Ocean Model:

Conservation laws for momentum, energy, salinity, mass and the equation of state

$$\begin{aligned}\frac{\partial \mathbf{v}}{\partial t} + (\mathbf{v} \cdot \nabla) \mathbf{v} + 2\boldsymbol{\omega} \times \mathbf{v} &= -\frac{\nabla M}{\rho} + \frac{\nabla \cdot \boldsymbol{\tau}}{\rho}, \\ \frac{\partial(\Delta h T)}{\partial t} + \nabla \cdot (\Delta h T \mathbf{v}) &= \nabla \cdot (\nu \Delta h \nabla T) + F^T, \\ \frac{\partial(\Delta h S)}{\partial t} + \nabla \cdot (\Delta h S \mathbf{v}) &= \nabla \cdot (\nu \Delta h \nabla S) + F^S, \\ \frac{\partial}{\partial t}(\Delta h) + \nabla \cdot (\mathbf{v} \Delta h) &= 0, \\ \rho &= \rho(T, S, P),\end{aligned}$$

The model considers 40 isopycnal layers and the outputs are provided over a uniform 2250×4500 resolution grid.



Quasi-geostrophic double-gyre model:

stratified potential vorticity equations
with meridional gradient of planetary
vorticity, lateral viscosity & bottom
friction and wind forcing

$$\begin{aligned}\partial_t q_i + J(\psi_i, q_i) &= \delta_{1i} F_w - \delta_{i3} \frac{a_v}{H_3^2} \nabla^2 \psi_i + a_h \nabla^4 \psi_i \\ q_i &= \nabla^2 \psi_i + \beta y - (1 - \delta_{i1}) S_{i1} (\psi_i - \psi_{i-1}) - (1 - \delta_{i3}) S_{i2} (\psi_i - \psi_{i+1})\end{aligned}$$

(Karabasov et al. 2009, Shevchenko & Berloff 2015)

boundary conditions

$$\partial_{\mathbf{n}\mathbf{n}} \psi_i - \alpha^{-1} \partial_{\mathbf{n}} \psi_i = 0$$
$$i = 1, 2, 3$$

numerical
solution

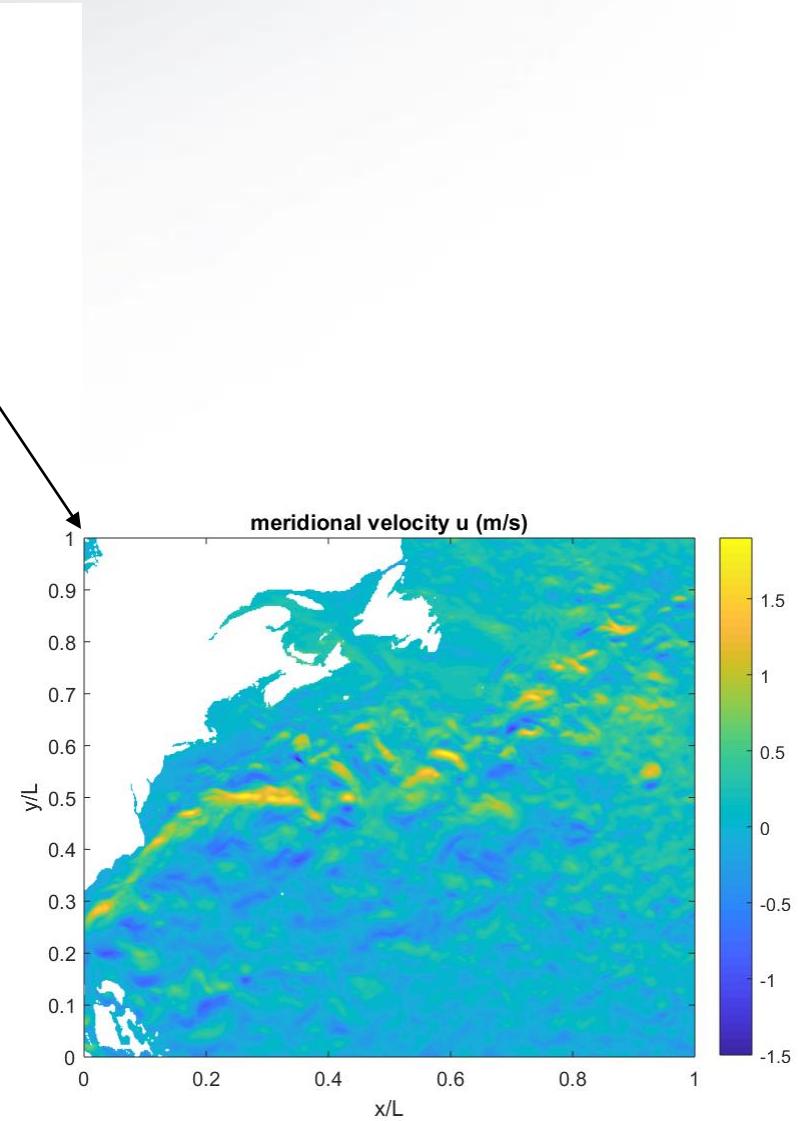
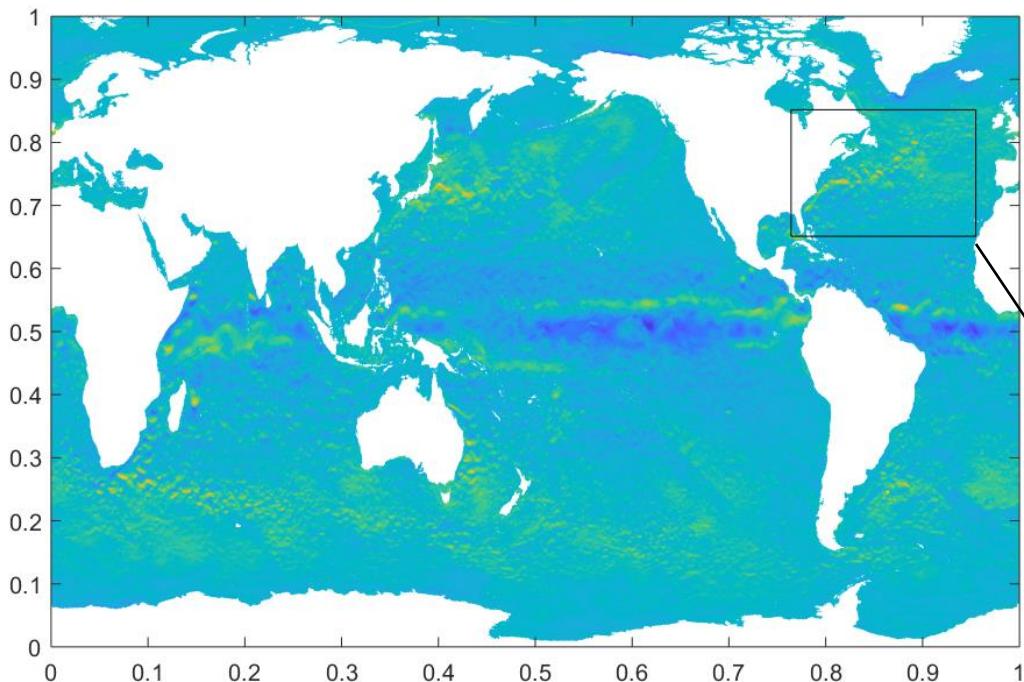
CABARET: Compact Accurately Boundary-Adjusting high-REsolution
Technique (Karabasov & Goloviznin 2009)

package

PEQUOD: Parallel Quasi-Geostrophic Model, (Maddison, Berloff &
Karabasov 2014)

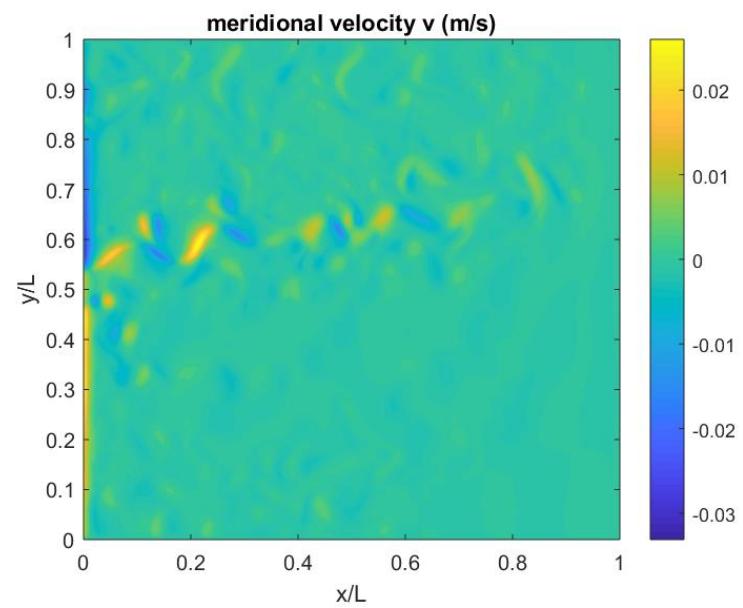
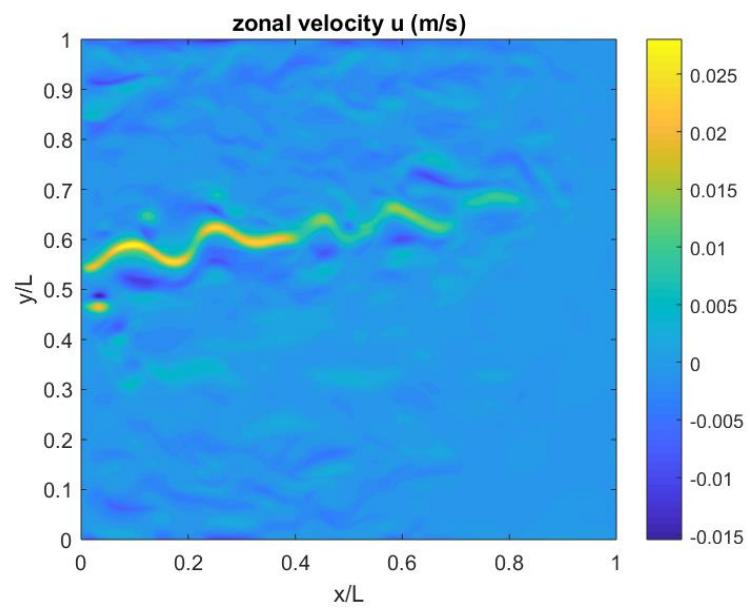


HYCOM outputs

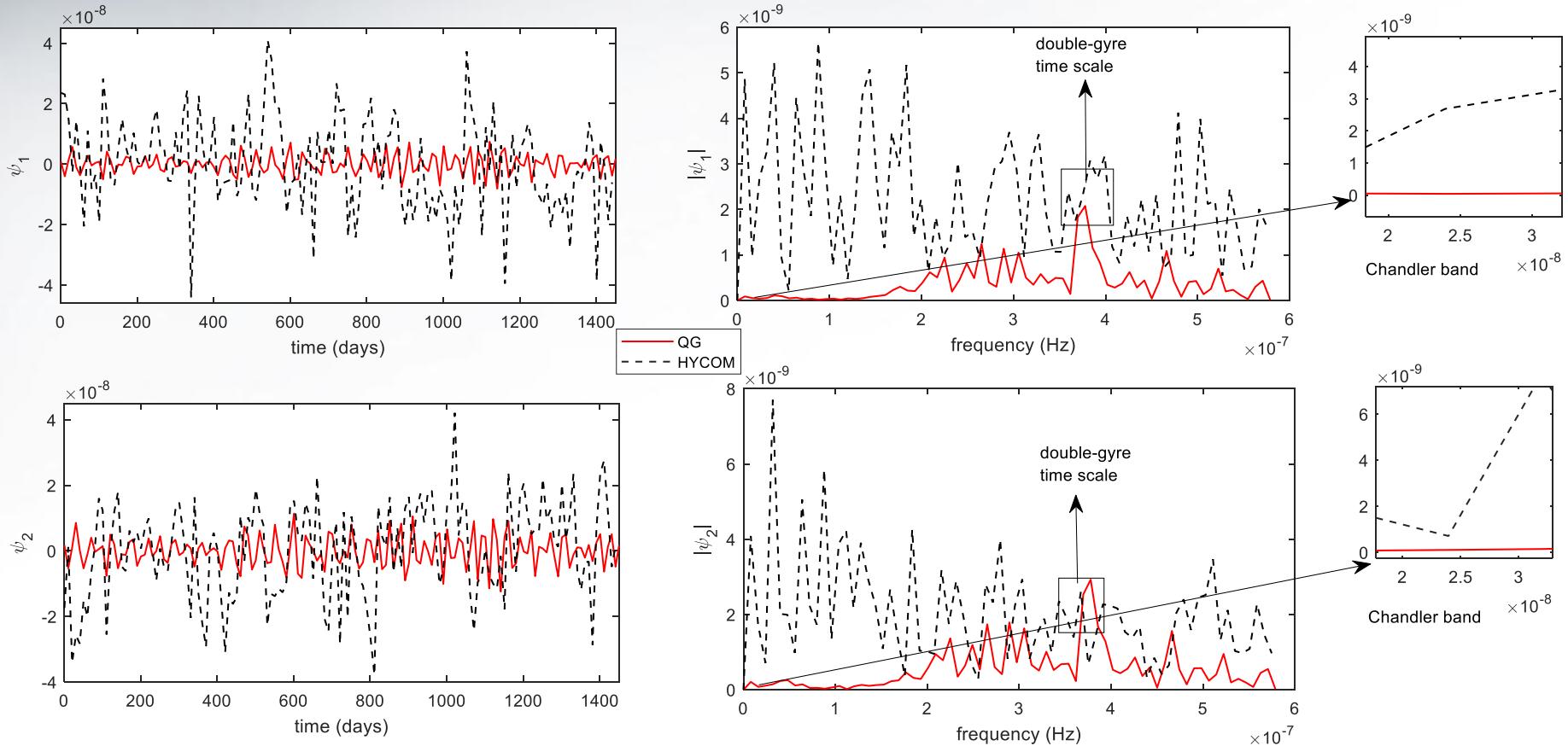




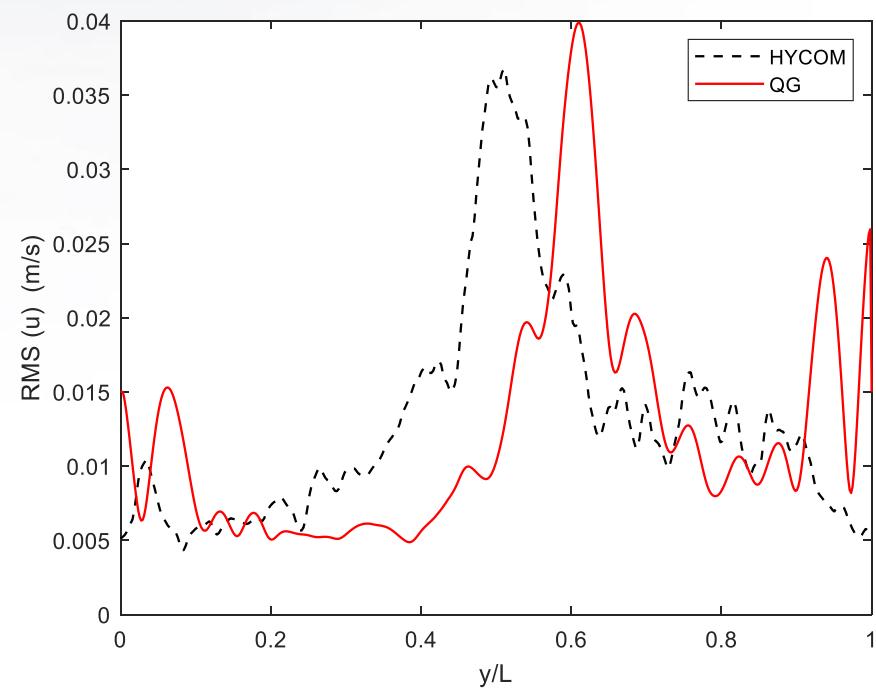
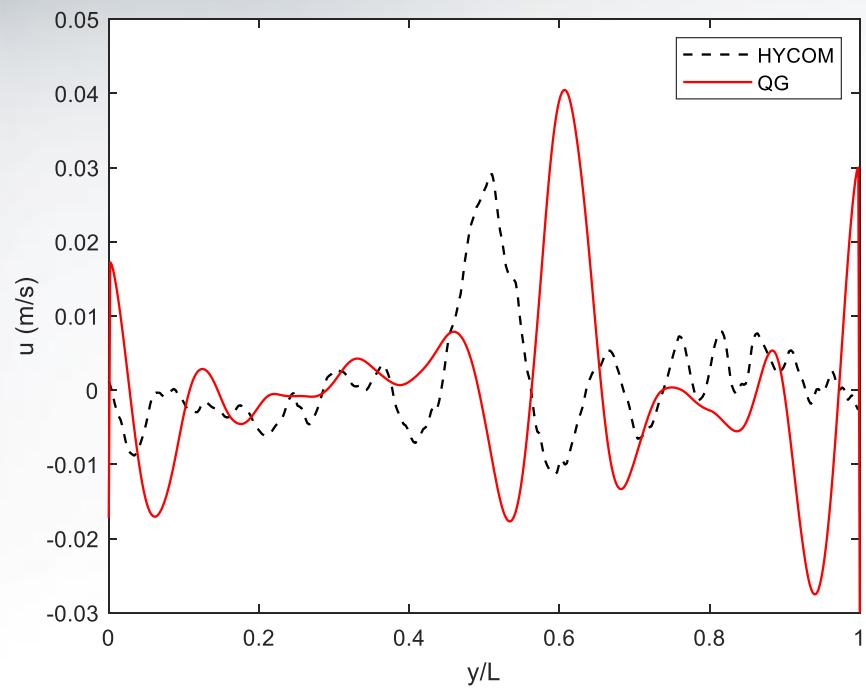
QG outputs



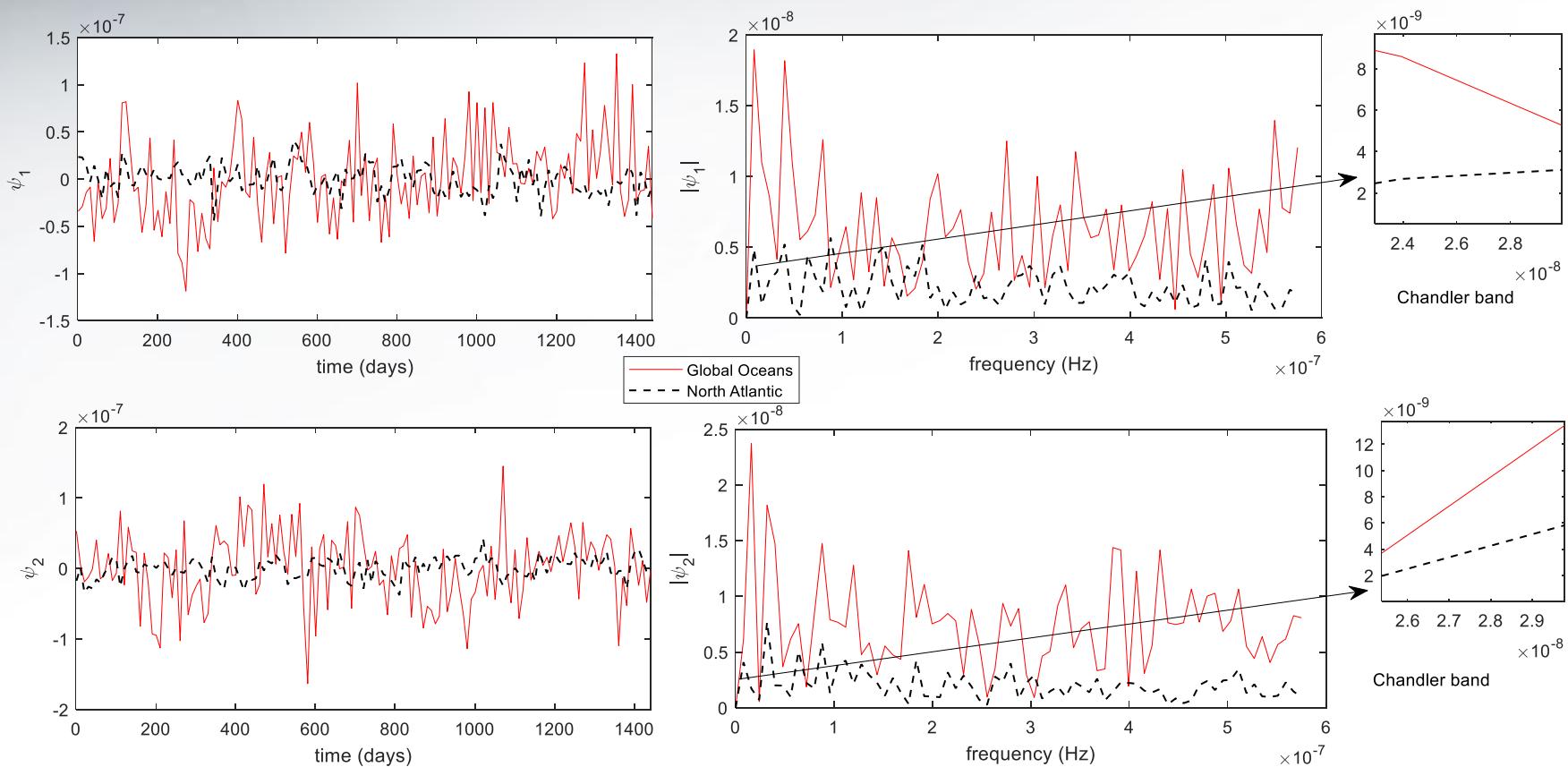
Comparison of Chandler wobble excitation functions using HYCOM & QG velocity fields



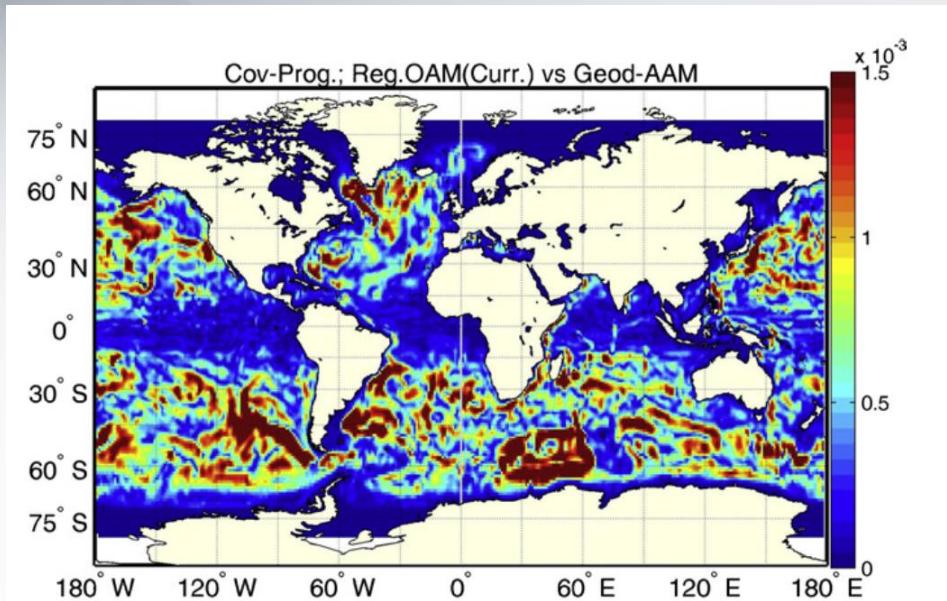
Comparison of the meanflow and RMS profiles in HYCOM & QG



Chandler wobble excitation functions: Global oceans vs. the North Atlantic (motion term)



North Atlantic is not the largest contributor in Chandler wobble excitation



Nastula et al. 2012

Excitation Process	Power, mas ²
Observed	1.26
Pacific Ocean	
χ_{uv}^{PO}	0.08(6.4%)
χ_p^{PO}	0.46(36.5%)
χ_{uv+p}^{PO}	0.28(22.2%)
Indian Ocean	
χ_{uv}^{IO}	0.18(14.3%)
χ_p^{IO}	0.18(14.3%)
χ_{uv+p}^{IO}	0.16(12.7%)
Atlantic Ocean	
χ_{uv}^{AO}	0.02(1.6%)
χ_p^{AO}	0.15(11.9%)
χ_{uv+p}^{AO}	0.09(7.1%)

Ma et al. 2009

(Naghibi et al., 2017)



Future work:

- Calculating Chandler wobble excitation over a longer period of time using HYCOM & QG outputs
- Comparing the predicted excitation with geodetic observation of the Chandler wobble after subtraction of mass terms estimated from geophysical GRACE satellite gravimetry data



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Thank you